

## Increased complexity (Example of landfall)

- Axi-symmetric models
  - (Ooyama, 1969, Rosenthal, 1971)
- Arbitrarily cut evaporation CE=0
- > 3-D (Tuleya and Kurihara, 1978)
- Steering current ---move b.c. through domain
- Predict sfc temperature (requires sfc fluxes including radiation)
- Predict moisture at surface
- Model land surface more completely

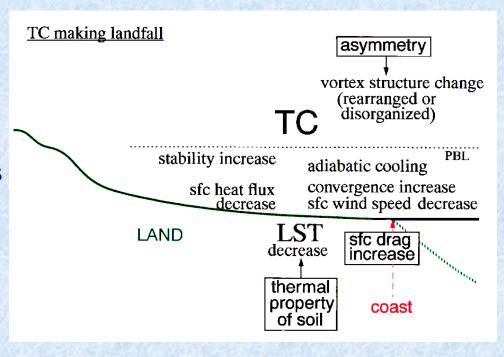


TABLE 8.—Experiments that examine the relative importance of air-sea exchanges of sensible heat, latent heat, and momentum during the mature stage of the model cyclone. The initial data are from hour \$88 of experiment \$35.

Experiment	$C_D$	Ca	CL	
835	eg (3)	eq (3)	eq (8)	
Q13	do.	0	do.	
Q14	do.	eq (8)	Ď	
Q15	0	do.	eg (3)	
Q16	ed (3)	2×eq (3)	2Xeq (8)	

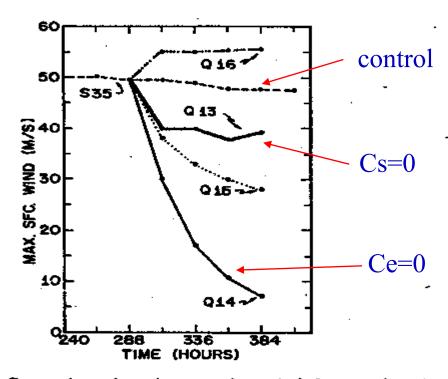
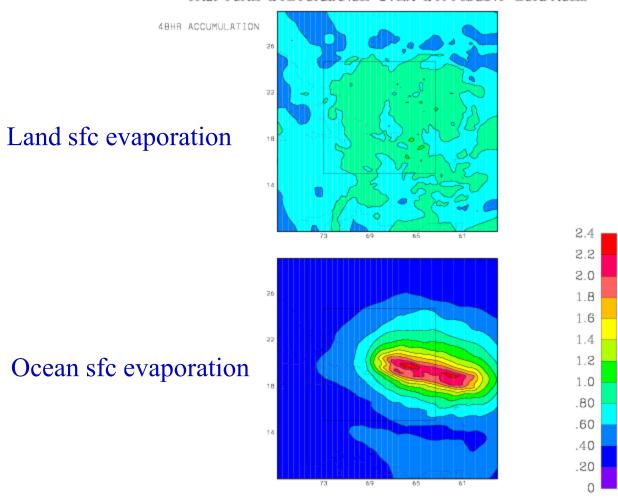


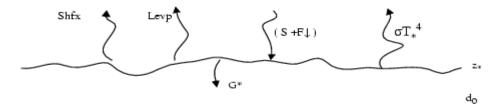
FIGURE 12.—Comparison of maximum surface winds for experiments that study the relative importance of air-sea exchanges of sensible heat, latent heat, and momentum during the mature stage of the model cyclone. Initial data are taken from hour 288 of experiment S35. The experiments compared with the control are Q13  $(C_0=0)$ , Q14  $(C_L=0)$ , Q15  $(C_D=0)$ , and Q16  $(C_0=C_L=2C_D)$ . See table 8 for details.

#### TROPICAL DISTURBANCE OVER DIFFERENT SURFACES



#### HWRF & GFDL use slab ground surface layer

#### Land Surface Temperature Prediction in the GFDL Model



#### 1. Formulation

Assume surface energy balance

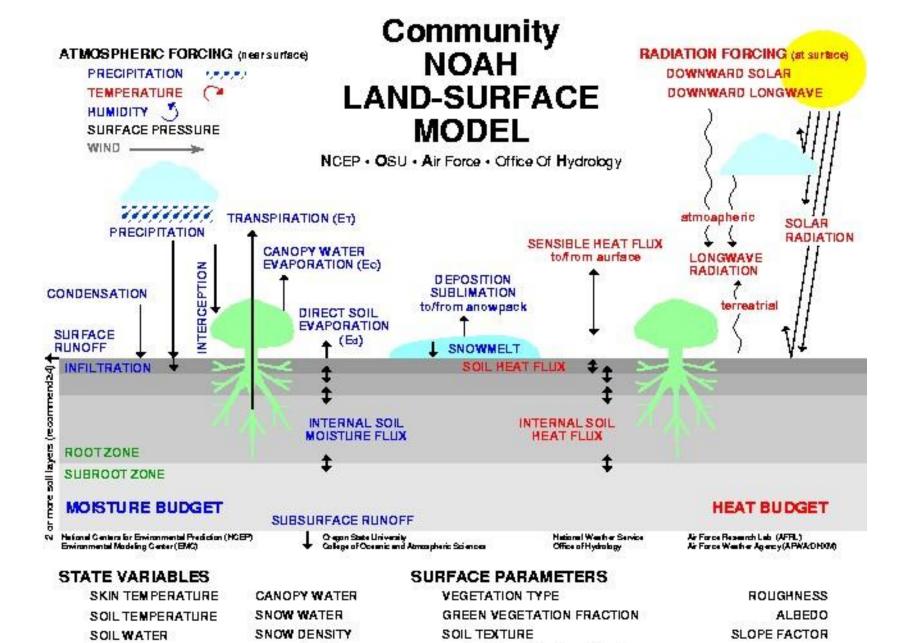
$$\sigma T_*^4 + Shfx + Levp \cdot (S + F_{\downarrow}) = G_*$$
 (1)

Assume diurnal temperature variation extends to depth,  $d_0$ , where the ground heat flux  $G(z) \Rightarrow 0$ . In this layer from  $z_*$  to  $d_0$  the vertical variation of heat flux may be related to the temperature tendency by  $\partial G/\partial z \propto \rho_* c_* \partial T/\partial t$  (1.1). Classic 'equilibrium' approach is to assume  $G_* = 0$ . Other studies have related  $G_*$  to be  $f(T_*$  or  $\partial T_*/\partial t)$  for diagnosis or prediction of  $T_*$ . Some relevant references are in Deardorff (1978, JGR), Bhumralkar (1975, JAM) and Blackadar (1976). Here we integrate (1.1) from the surface to  $d_0$ :

where 
$$\partial T/\partial t \equiv [\partial T/\partial t]_* e^{-z/d}$$
  
and d, the damping depth,  $\equiv (\nabla k_z/\pi)^{1/2}$  is defined in Sellers  
and  $\tau$  is the period of forcing (24hrs)  
 $G_* = \rho_s c_s d \partial T_*/\partial t$  (1 - ~ 0) for  $d_0 >> d$   
 $G_* = \rho_s c_s d \partial T_*/\partial t$  (2)

The temperature tendency,  $\partial T_w/\partial t$ , can then be calculated from (2) and (1):

$$\partial T_*/\partial t = (\sigma' T_*^4 + \mathbf{Shfx} + \mathbf{Levp} \cdot (\mathbf{S} + \mathbf{F}\downarrow)) / (\rho_s c_s d)$$
 (3)



ftp://ftp.ncep.noaa.gov/pub/gcp/ldas/noahlsm/

SOILICE

## EMC transitioning to Noah LSM & The EMC Stream Routing Scheme:

- 1. Computes the concentration time for runoff reaching the outlet of a grid box and the transport of water in the channel system
- 2. Water can leave the grid cell through (at least) one of the eight directions
- 3. The runoff transport process is linear and time invariant
- 4. The causality and the impulse response functions are nonnegative

## **Experimental Design**

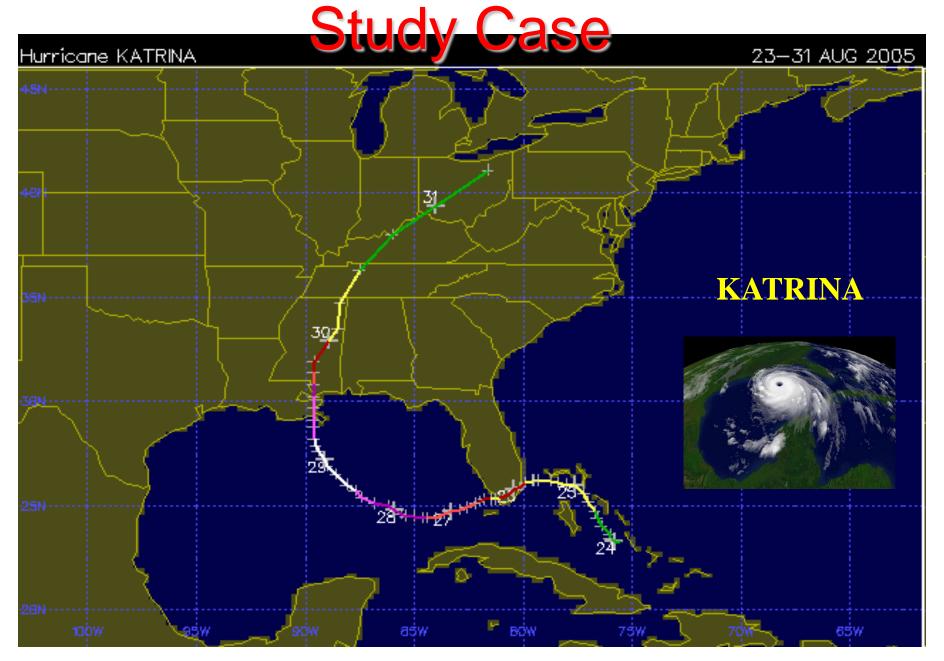
Yihua Wu (EMC)

Runs	Physics Options Over Land			Physics Options Over Ocean			
	LSM	Surface Layer Scheme	PBL Scheme	Surface Layer Scheme	PBL Scheme		
	Noah	GFDL	GFS	GFDL	GFS		
N893.WET	The initial soil moisture for US is from GFS soil moisture						
N893.DRY	The initial soil moisture for US is the combination of GFS soil moisture and NAM soil moisture						

### **Preliminary Results**

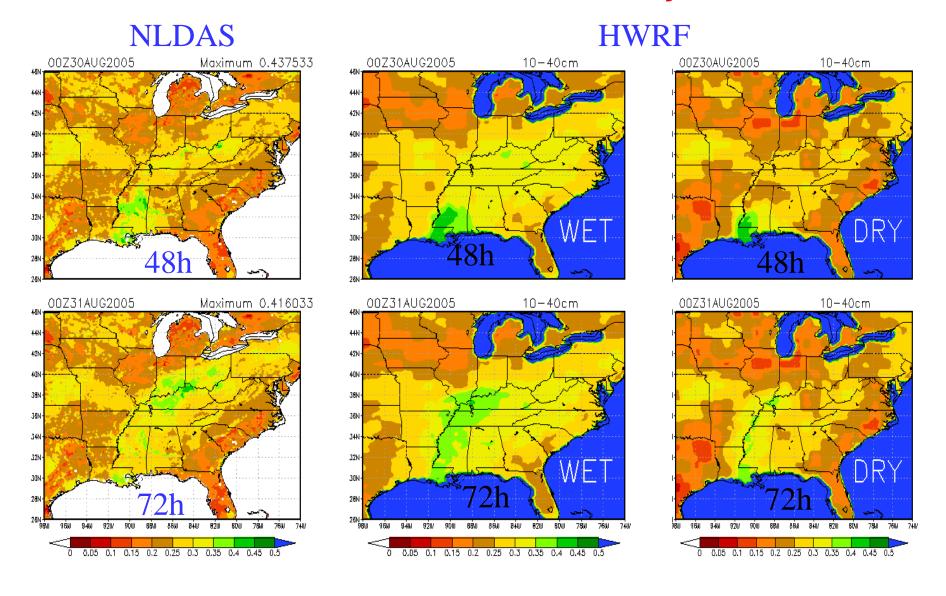
~dozen cases

- Track forecasts not impacted by Noah Ism
- Intensity change appears relatively minor
- Some impact on rainfall forecasts
- Runfoff and stream flow impacted by initial soil moisture
- Need to run extensive tests

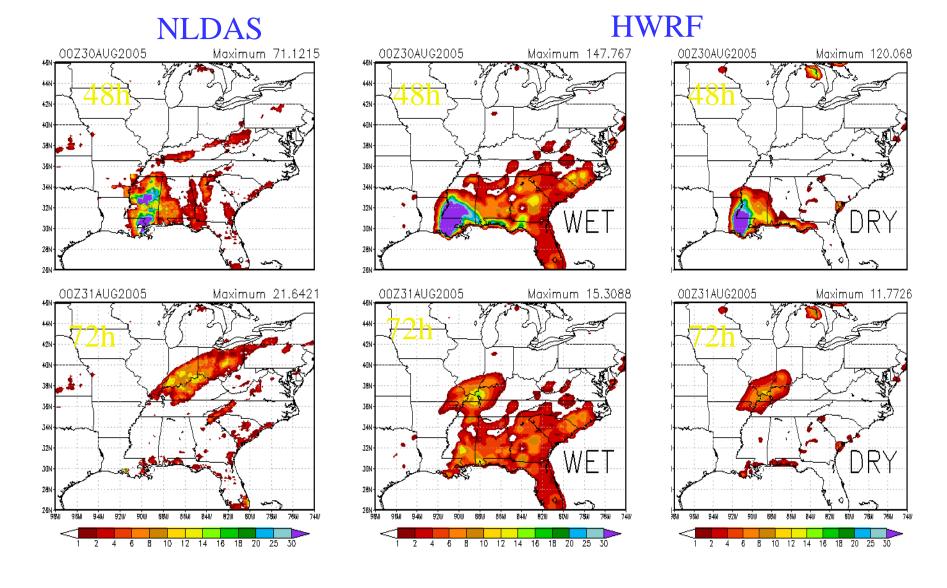


All runs for Katrina started at 00z Aug 28, 2005

### **Forecasted Soil Moisture for Layer 2**

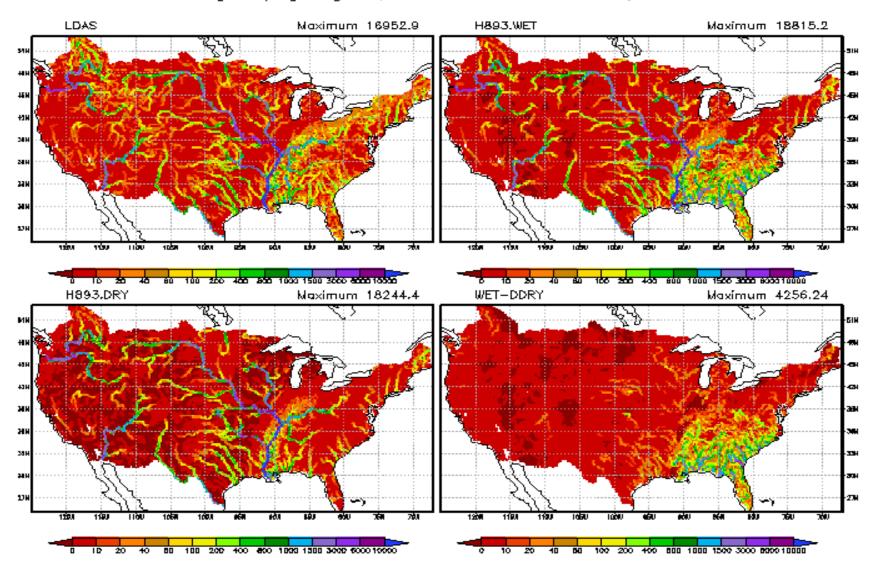


### 12 Hour Total Runoff (mm)

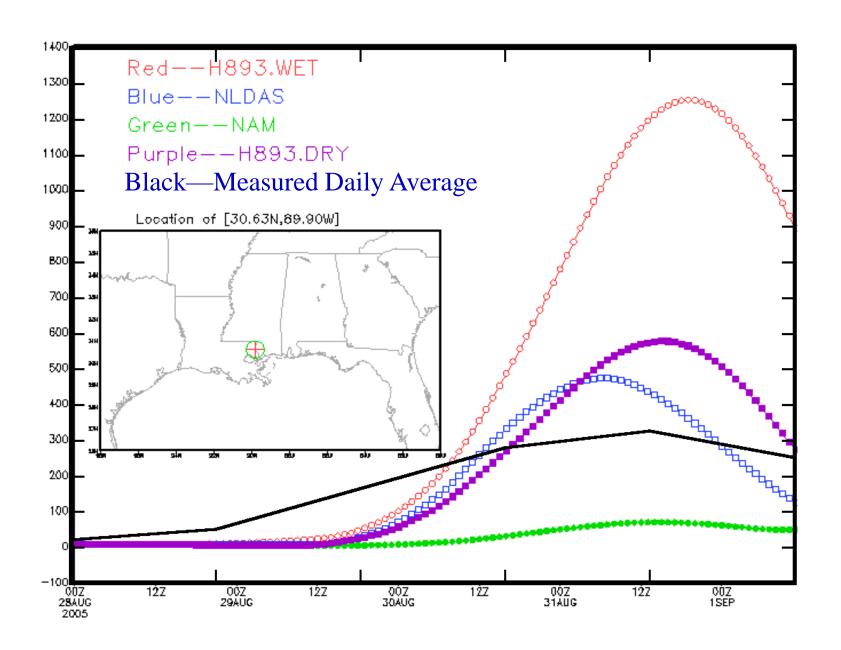


#### Forecasted Stream Flow (m<sup>3</sup> s<sup>-1</sup>)

Sreamflow [m^3/s] Aug 28, 2005 00Z: 95 H FCST, 23Z31AUG2005



#### Stream flow At 30.63N and 89.90W



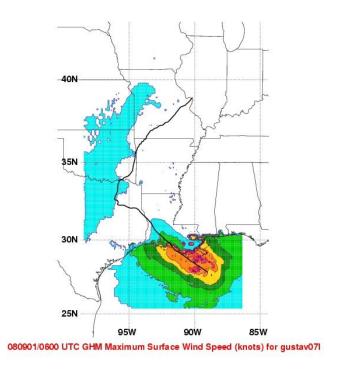
## Hurricane-specific model products

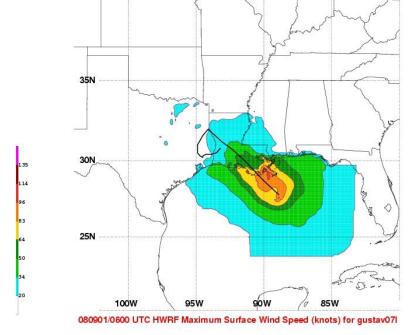
- Storm track & intensity (i.e. "tracker")
- Low level wind swaths ... Meow...

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1st analysis?? ... Ted Fujita on Camille (1969)
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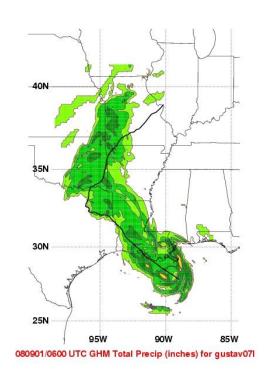
- Rainfall swaths
- Be aware of analysis methods used in products above

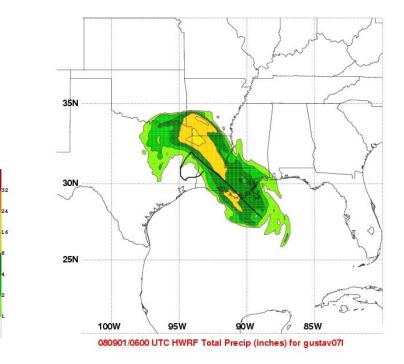
## Low level wind swaths GFDL HWRF





## Storm total precip GFDL HWRF





### Inland decay model

(Kaplan and DeMaria 1995, 2001 & 2006)

### (a more simpler approach)

- Assumes exponential decay
- Uses climatology
- Can be used as a baseline for 3-d models
- Can be driven by a variety of forecasts

### Background

An empirical decay model (Kaplan and DeMaria 1995, 2001) has been used to predict the decrease in wind speed of landfalling tropical cyclones. The model assumes that a cyclone's maximum winds decrease exponentially with time after landfall to a non-zero background wind speed using:

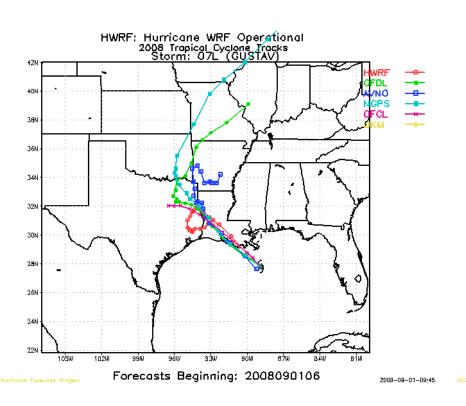
$$V_t = V_b + (RV_0 - V_b)e^{-\alpha t}$$

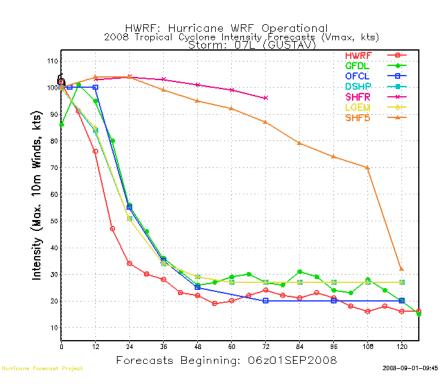
where  $V_t$  = the maximum wind at some time t after landfall,  $V_0$  is the landfall wind speed,  $V_b$  is the background wind speed and  $\alpha$  is the decay constant.

DeMaria et al. (2006) recently developed a revised version of the original Kaplan and DeMaria decay model that improves the prediction for storms that cross islands and peninsulas. The new version decreases the rate of decay of landfalling storms according to the fractional area of the storm that is over land (F<sub>m</sub>) during any given time and is given by:

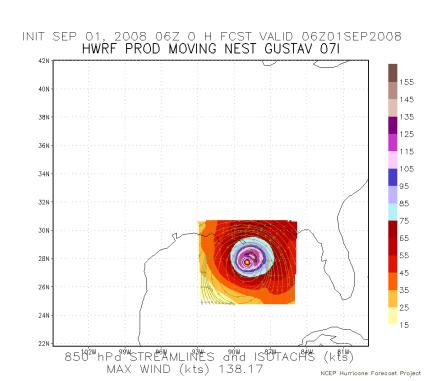
$$V^{t+1} = V_b + (V^t - V_b)e^{-F_m \cot}$$

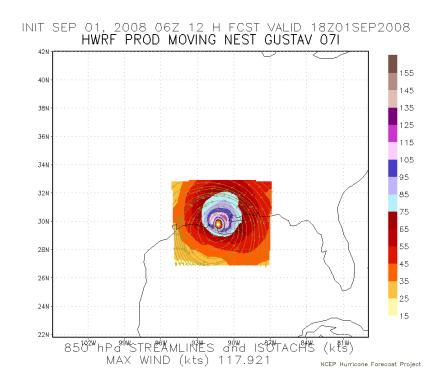
## Operation models of Gustav at landfall



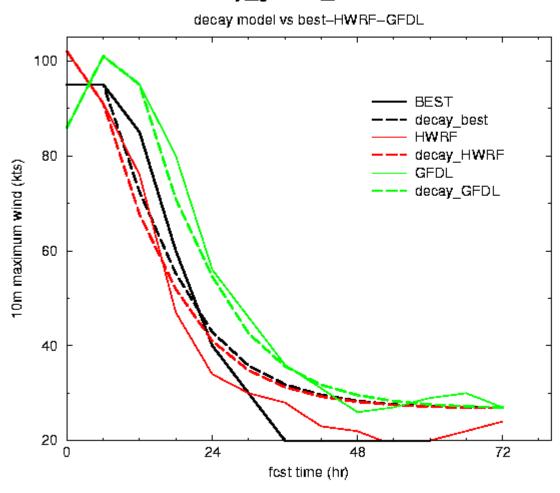


### **HWRF** forecast

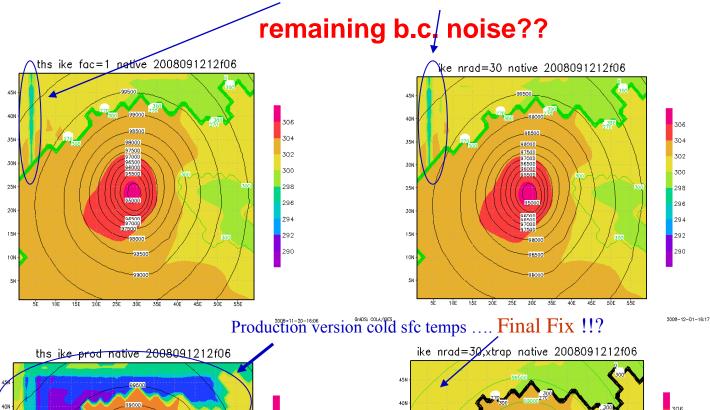


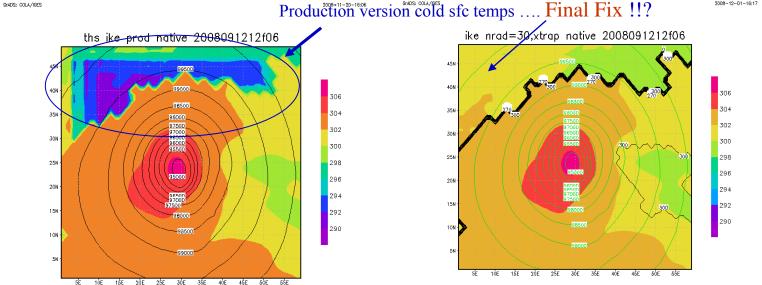


#### decay\_gustav\_08090106



### **HWRF THS at f=6hrs IKE**





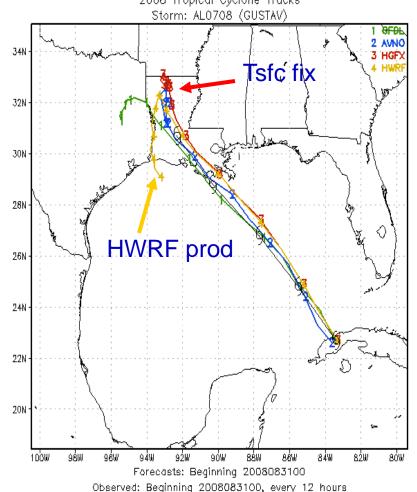
GrAUS; COLA/RES 2008-10-07-16:12

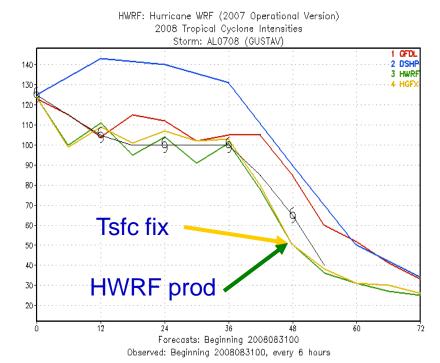
RHADS: COLLAZIONS

2008-12-19-16:21

## Gustav 083100 Impact of Tsfc fix: (improved track, same intensity)

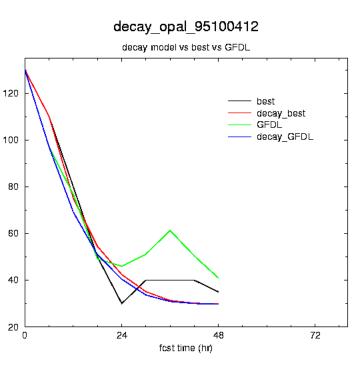
HWRF (2007 Operational Version) Coupled Model Forecasts
2008 Tropical Cyclone Tracks

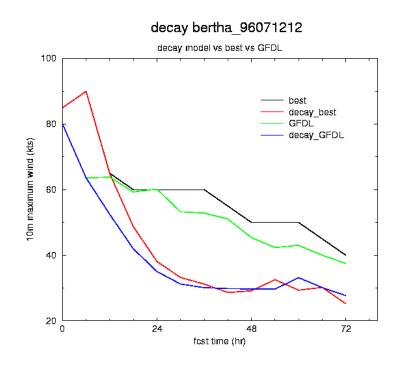




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## more examples of GFDL and inland decay model



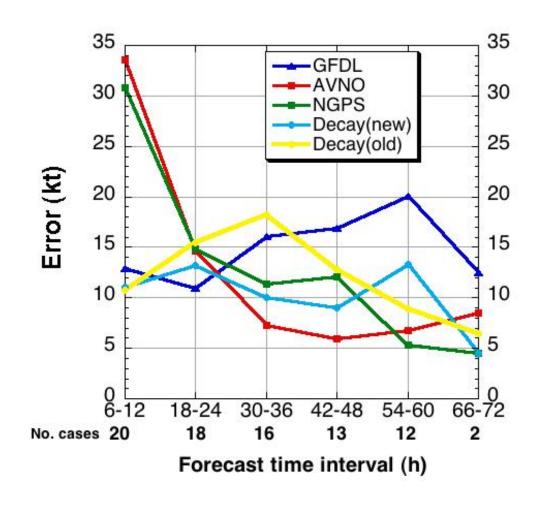


## More systematic verification of decay over land

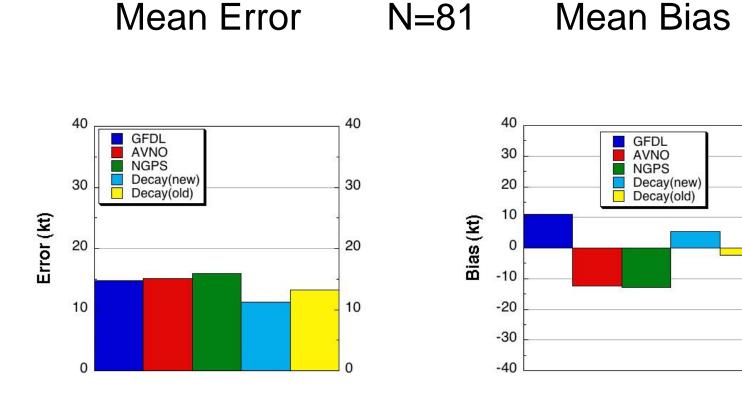
John Kaplan

- Compared inland decay model to NWP models
- Need to compare with HWRF and more cases
- Nagging problem of observations??

#### Absolute error in the model forecasted maximum wind



Sample mean absolute error and mean bias of the model maximum wind forecasts for all time intervals (0-72 h)



40

30

20

10

0

-10

-20

-30

-40

### **Evaluation of 2-d fields**

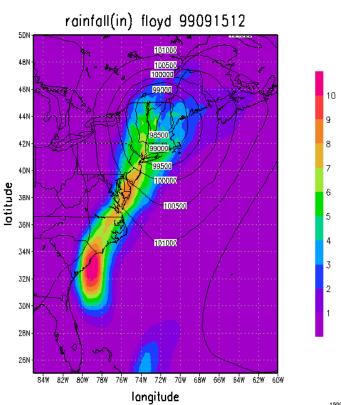
### Rainfall & Wind

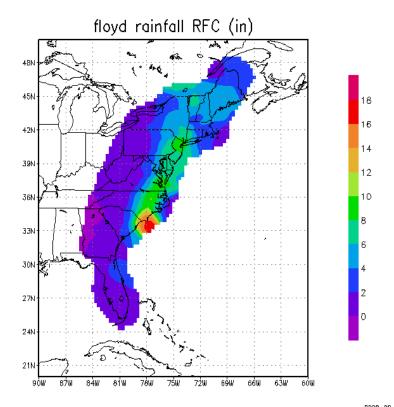
Evaluation metrics

Absolute mean error, rms, bias&threat, pdf, taylor diagram??

Spatial and temporal scale??

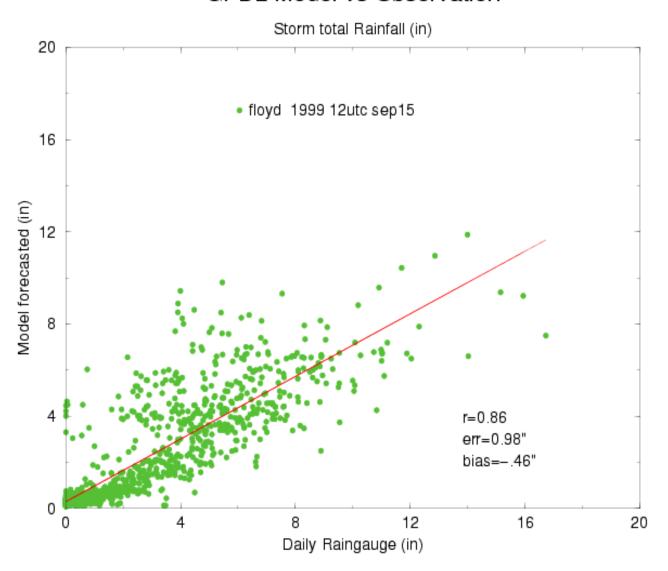
## Floyd 091512 model vs. daily gauges





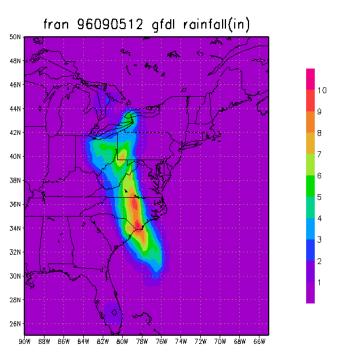
6r/MS: COLA/IBES 1998\_-(B-M/US: COLA/IBES 1998\_-(B-M/US: COLA/IBES 2002-02-12-16:07

#### GFDL Model vs Observation

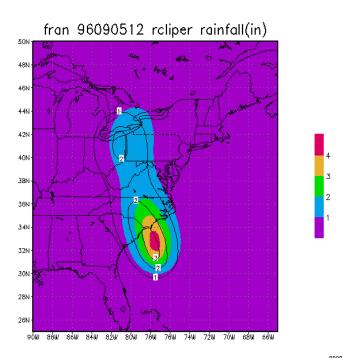


### Model vs Rainfall Cliper

#### Model



#### Best vs Model track Reliper



Grads: COLA/IGES 2001-07-23-10:41 Grads: COLA/IGES 2001-07-23-10:41 Grads: COLA/IGES 2002-03-04-16:35

# Evaluation of 2-d fields wind, rainfall, etc issues

- Evaluate from nhc 'vitals' file restricted to incremental quadrants vitals file data valid over land??
- Evaluate at observations location (e.g. Tuleya et al, 2007)
   non-homogeneity of observation type and coverage
- Evaluate on homogeneous analysis grid

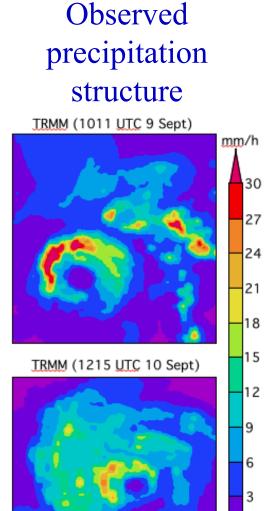
(e.g Marchok et al, 2007)

blends in all data types normalize both observations and model data to standard reference values..i.e. standard open terrain

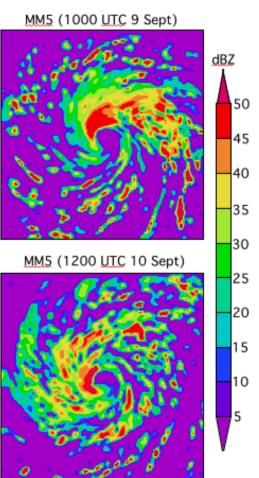
### How to assess small scales ??

- High resolution fields appear more realistic but....
- Small scale fields reduce 'skill' if out of phase with observations (e.g. GFS vs GFDL and ETA models rainfall)
- Need to transition to ensemble or probabilistic approach

### **Simulation of Hurricane Erin (2001)**

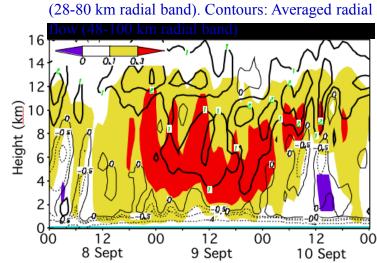


## Simulated precipitation structure



- Observed and simulated precipitation show shift of rainfall to western side from Sept. 9 to 10 (left).
- •This shift corresponds to a weakening of the eyewall vertical motions and switch from midlevel outflow to inflow (below).

Colors: Azimuthally averaged vertical motion



L. Wu and S. Braun, NASA/GSFC

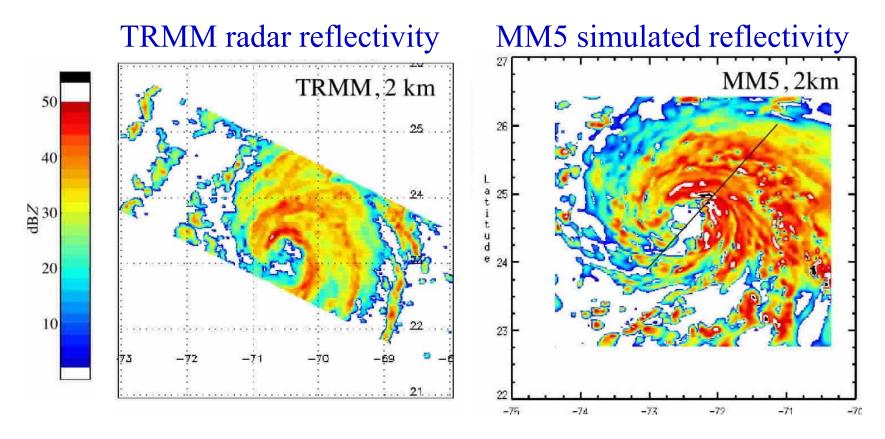
### SUMMARY

### Starting to evaluate 2-D fields

- More than track and intensity
- Realistic patterns may not lead directly to increase skill
- Advocate using parametric models as baseline and use directly with model data track & intensity to separate out track & intensity error effects
- Comments and suggestions please!!!

## Validation of a Hurricane Bonnie Simulation using TRMM Precipitation Radar

- MM5 simulation with innermost grid of 2-km grid spacing
- Model reproduces asymmetry and multiple convective rainbands



#### S. Braun, NASA/GSFC